

## Thicker shells compensate extensive dissolution in brachiopods under future ocean acidification

Emma Cross, Elizabeth M. Harper, and Lloyd S. Peck

*Environ. Sci. Technol.*, **Just Accepted Manuscript** • DOI: 10.1021/acs.est.9b00714 • Publication Date (Web): 29 Mar 2019

Downloaded from <http://pubs.acs.org> on April 3, 2019

### Just Accepted

"Just Accepted" manuscripts have been peer-reviewed and accepted for publication. They are posted online prior to technical editing, formatting for publication and author proofing. The American Chemical Society provides "Just Accepted" as a service to the research community to expedite the dissemination of scientific material as soon as possible after acceptance. "Just Accepted" manuscripts appear in full in PDF format accompanied by an HTML abstract. "Just Accepted" manuscripts have been fully peer reviewed, but should not be considered the official version of record. They are citable by the Digital Object Identifier (DOI®). "Just Accepted" is an optional service offered to authors. Therefore, the "Just Accepted" Web site may not include all articles that will be published in the journal. After a manuscript is technically edited and formatted, it will be removed from the "Just Accepted" Web site and published as an ASAP article. Note that technical editing may introduce minor changes to the manuscript text and/or graphics which could affect content, and all legal disclaimers and ethical guidelines that apply to the journal pertain. ACS cannot be held responsible for errors or consequences arising from the use of information contained in these "Just Accepted" manuscripts.

# Thicker shells compensate extensive dissolution in brachiopods under future ocean acidification

Emma L. Cross<sup>†, ‡, \*</sup>, Elizabeth M. Harper<sup>†</sup>, Lloyd S. Peck<sup>‡</sup>

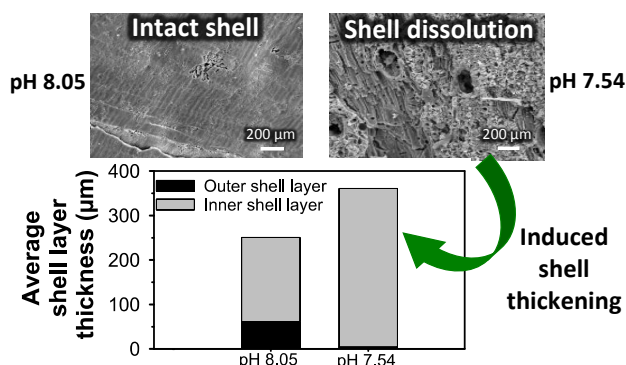
<sup>†</sup>Department of Earth Sciences, University of Cambridge, Downing Street,  
Cambridge, CB2 3EQ, UK.

<sup>‡</sup>British Antarctic Survey, Natural Environment Research Council, High Cross,  
Madingley Road, Cambridge, CB3 0ET, UK

KEYWORDS: climate change, global warming, shell dissolution, shell thickness, compensatory mechanisms, phenotypic plasticity, *Calloria inconspicua*, *Liothyrella uva*, Terebratulide

## ABSTRACT

Organisms with long generation times require phenotypic plasticity to survive in changing environments until genetic adaptation can be achieved. Marine calcifiers are



particularly vulnerable to ocean acidification due to dissolution and a reduction in shell-building carbonate ions. Long-term experiments assess organisms' abilities to acclimatise or even adapt to environmental change. Here we present an unexpected compensatory response to extensive shell dissolution in a highly calcium-carbonate-dependent organism after long-term culture in predicted end-century acidification and warming conditions. Substantial shell dissolution with decreasing pH posed a threat to both a polar (*Liothyrella uva*) and a temperate (*Calloria inconspicua*) brachiopod after 7 months and 3 months exposure, respectively, with more extensive dissolution in the polar species. This impact was reflected in decreased outer primary layer thickness in the polar brachiopod. A compensatory response of increasing inner secondary layer thickness, and thereby producing a thicker shell was exhibited by the polar species. Less extensive dissolution in the temperate brachiopod did not affect shell thickness. Increased temperature did not impact shell dissolution or thickness. Brachiopod ability to produce a thicker shell when extensive shell dissolution occurs suggests this marine calcifier has great plasticity in

35 calcification providing insights into how similar species might cope under future  
36 environmental change.

## INTRODUCTION

Changing environments pose serious risks to organisms that cannot shift their geographic range, physiologically acclimatise or genetically adapt<sup>1</sup>. Current understanding of the biological impacts of ocean acidification and warming is largely based on short- (days) to medium-term (weeks) laboratory and field experiments that have revealed mixed responses in many species<sup>2-5</sup>. More recently, however, there has been an increase in long-term (many months to years) studies that demonstrate surprising capacities of marine organisms to acclimate<sup>6-10</sup>, or even adapt in organisms with short generation times<sup>11-13</sup> to decreased pH and increased temperature. Compensatory mechanisms could be paramount to maintain overall performance of organisms that have limited capacities to alter their geographic range under future changed conditions and subsequently sustain their key ecological functions in our oceans<sup>14</sup>.

Marine calcifiers are considered the most vulnerable organisms to ocean acidification due to the combination of dissolution and the reduction in carbonate ions making shell production more difficult and energetically expensive<sup>2,15,16</sup>. The Southern Ocean has naturally low carbonate ion saturation levels compared to temperate and tropical regions due to carbon dioxide being more soluble in cold water<sup>17</sup>. Acid-base coefficients are also more sensitive in cold temperatures making this high latitude region a forerunner of biological ocean acidification impacts for other oceans<sup>18</sup>. The external skeleton is crucial for protecting animal tissue in shell-bearing organisms against predation, infection and loss of bodily fluids<sup>19,20</sup>. Any

negative impacts to shell integrity, therefore, could compromise its protective function and potentially prove fatal. Shell integrity may be affected by erosion from natural scour or attack from shell-boring organisms as well as dissolution. The calcified shell of all shell-bearing organisms is protected by an outer organic layer, the periostracum<sup>21,22</sup>. Abrasion of this protective layer and subsequently inner shell layers naturally occurs through abrasion from suspended inorganic particulate material, the movement of individuals against each other, and with other calcified biota or substrata. Shell dissolution also poses a threat depending on the solubility of the biomineral, the chemical characteristics of the surrounding seawater and metabolic by-products released by the adhering biofilm<sup>23,24</sup>. Predicted environmental conditions for 2100 will shift surface seawater carbonate chemistry to favour  $\text{CaCO}_3$  dissolution, which could exacerbate the loss of shell integrity of marine calcifiers.

Compensatory mechanisms may counteract deleterious ocean acidification and warming effects on organisms. For these to succeed, the compensatory mechanism must occur at a faster rate than that of the deleterious effect to provide successful protection. Phenotypic plasticity of shell morphology has been reported in shelled organisms in response to the presence of predators<sup>25</sup> and changing environmental conditions<sup>14,26-28</sup>. These include shell thickening, production of a more rotund shell and increased shell growth rates through plasticity in producing different calcium carbonate polymorphs<sup>14,25-32</sup>. Production of a thicker periostracum could also withstand more wear and deter dissolution<sup>33</sup>. Periostracum loss or shell dissolution at the external surface far away from the secretory tissue cannot be directly repaired

by the organism. Compensatory mechanisms such as induced thickening, however, could counteract this potentially fatal effect of ocean acidification.

Brachiopods are one of the most calcium-carbonate-dependent groups of marine animals because their calcareous skeleton and other support structures make up > 90% of their dry mass<sup>34,35</sup>. Rhynchonelliform brachiopods possess a low-magnesium calcite shell consisting of the periostracum underlain by two biomineralised inner layers; the thin nanocrystalline primary layer and the generally much thicker fibrous secondary layer<sup>36,37</sup>. In previously published work we showed that shell growth rates of *L. uva* Broderip, 1833 (which we refer to as “polar brachiopod”) and *C. inconspicua* Sowerby, 1846 (which we refer to as “temperate brachiopod”) were not impacted by predicted end-century seawater pH’s<sup>6,7</sup>. Another study demonstrated increased dissolution in the polar brachiopod in pH 7.4 conditions after 14 days<sup>38</sup>, however, empty dried valves were used so the brachiopods ability to compensate shell dissolution remains unknown. This study, therefore, investigated dissolution effects and potential compensatory mechanisms of a polar and a temperate brachiopod living under acidified and warming conditions. Specifically, the extent of dissolution and thickness of whole valves and individual shell layers were assessed under predicted end-century pH levels in both brachiopods and also under increased temperature in the polar brachiopod.

## MATERIALS AND METHODS

**Sampling collection.** Specimens of the polar brachiopod were hand collected by SCUBA divers from Trolval Island, Ryder Bay, Antarctica (67° 35.44' S, 68° 12.44' W) at 15-25 m depth in May 2012. Environmental conditions in Ryder Bay at 15-25 m depth consist of seawater temperatures that range from -1.8 to +1.5°C, however, temperatures rarely exceed +1.0°C and salinity is 33.0-34.0<sup>39</sup> and the pH range is 8.04-8.10<sup>40</sup>. Brachiopods were kept in recirculating aquaria (0.0 ± 0.5°C) whilst being transported by ship back to the British Antarctic Survey, Cambridge, UK where the polar experiment was conducted.

Individuals of the temperate brachiopod were hand collected at low tide from under rocks in Portobello Bay, Otago Harbour, New Zealand (45° 82.00'S, 170° 70.00'E) in January 2013. Environmental conditions in Otago Harbour are surface seawater temperatures of 6.4-16.0°C<sup>41,42</sup>, pH range of 8.10-8.21 (K. Currie, pers. comm.) and salinity is 32.5-34.8<sup>42</sup>. Brachiopods were kept in seawater during the short transportation to Portobello Marine Laboratory, Otago Harbour, New Zealand where the temperate experiment was performed.

### Experimental Design.

**Polar experiment.** The polar experiment was conducted in a temperature-controlled recirculating CO<sub>2</sub> microcosm with four treatments<sup>6</sup>. Two were acidified treatments ("Moderate pH" – pH 7.75 ± 0.03 and "Low pH" – pH 7.54 ± 0.03) based



on the IPCC 'business-as-usual' scenario of the predicted end-century reduction of 0.3-0.5 pH units from the present day average of pH 8.1 in surface oceanic seawater by 2100<sup>43</sup> (Table 1). The third was a pH control where the seawater remained at ambient pH (pH  $8.05 \pm 0.03$ ). All these three treatments were maintained at 2°C throughout the experiment due to the concurrent 2°C increase in sea surface temperature (SST) expected to occur alongside these predicted decreased pH levels by the end of the century<sup>44</sup>. The fourth treatment was a temperature control which was held at the present-day average conditions for Ryder Bay<sup>45</sup> (SST: 0°C, pH:  $7.98 \pm 0.02$ ). The pH of the acidified treatments was controlled by intermittently bubbling CO<sub>2</sub> gas into a header tank. Seawater was then gravity fed into the experimental tanks<sup>6</sup>. The pH control treatment had a similar set up but without the pH manipulation system. The temperature control treatment was situated separately in the main BAS aquarium. Seawater temperature of all treatments was manipulated by controlling the air temperature in temperature-controlled laboratories.

Seawater temperatures (°C, Digital Testo 106) and pH<sub>NIST</sub> (Aquamedic pH controlled computer and electrode system) were monitored and recorded daily. Salinity (Tropical Marine Centre V2 Handheld refractometer), TCO<sub>2</sub> (mmol L<sup>-1</sup>; Ciba Corning TCO<sub>2</sub> Analyzer 965, Olympic Analytical. UK) and nutrient content (silicate and phosphate) of each treatment were measured weekly. Other carbonate system parameters, including the partial pressure of CO<sub>2</sub> (pCO<sub>2</sub>) and the saturation values for calcite ( $\Omega_C$ ) and aragonite ( $\Omega_A$ ), were modelled from applying TCO<sub>2</sub> and pH<sub>NIST</sub> data to the program CO2SYS<sup>46</sup> with refitted constants<sup>47,48</sup>. Brachiopods in each

147 treatment were fed weekly with microalgal concentrate of approximately  $331 \times 10^4$   
148 cells  $L^{-1}$ , which is within the natural seasonal range of phytoplankton cell abundance  
149 along the west Antarctic Peninsula ( $62\text{--}1150 \times 10^4$  cells  $L^{-1}$ )<sup>49,50</sup>.

150 *Temperate experiment.* The temperate experiment was conducted in a flow-  
151 through  $CO_2$  perturbation system with three treatments<sup>7</sup>. Two were acidified  
152 treatments ("Moderate pH" – pH  $7.79 \pm 0.06$  and "Low pH" – pH  $7.62 \pm 0.05$ ) and the  
153 third was a pH control ( $8.16 \pm 0.03$ ). The pH of the acidified treatments was lowered  
154 in header tanks by intermittently bubbling  $CO_2$  gas before being gravity fed into the  
155 replicate experimental tanks<sup>7</sup>. The pH control system had an identical set up except  
156 that it lacked  $CO_2$  injection, and air was injected into the header tank. Seawater  
157 temperature was not manipulated and was ambient for Otago Harbour.

158 Seawater temperatures ( $^{\circ}C$ , Digital Testo 106) and  $pH_{NIST}$  were measured three  
159 times a day and salinity (YSI data logger) was measured once a week. Dissolved  
160 inorganic carbon (DIC) and total alkalinity ( $A_T$ ) were analysed at the beginning,  
161 middle and end of the experiment by a Single Operator Multi-parameter Metabolic  
162 Analyser (SOMMA) and closed-cell potentiometric titration, respectively<sup>51</sup>. Other  
163 carbonate system parameters, including the partial pressure of  $CO_2$  ( $pCO_2$ ) and the  
164 saturation values for calcite ( $\Omega_C$ ) and aragonite ( $\Omega_A$ ) were calculated using CO2SYS<sup>46</sup>  
165 with  $CO_2$  equilibrium constants<sup>47,48,52</sup>. Brachiopods were fed three times a week with  
166 microalgal concentrate of approximately  $397 \times 10^4$  cells  $mL^{-1}$  of *Tetraselmis* spp.,  
167 which is within the natural summer range of phytoplankton cell abundance in Otago  
168 Harbour.

**Table 1.** Mean ( $\pm$ SD) seawater parameters during both the polar and temperate experiments.

Experiment	Treatment	pH <sub>NIST</sub>	Temperature (°C)	Salinity	$p\text{CO}_2$ ( $\mu\text{atm}$ )	$\Omega$ Calcite	$\Omega$ Aragonite
Polar	Temperature control	$7.98 \pm 0.02$	$-0.3 \pm 0.1$	$35 \pm 1$	$417 \pm 15$	$1.2 \pm 0.1$	$0.8 \pm 0.1$
	pH control	$8.05 \pm 0.03$	$1.7 \pm 0.3$	$35 \pm 1$	$365 \pm 67$	$1.5 \pm 0.2$	$0.9 \pm 0.1$
	Moderate pH	$7.75 \pm 0.03$	$1.9 \pm 0.4$	$35 \pm 1$	$725 \pm 133$	$0.8 \pm 0.1$	$0.5 \pm 0.1$
	Low pH	$7.54 \pm 0.03$	$2.2 \pm 0.4$	$35 \pm 1$	$1221 \pm 179$	$0.5 \pm 0.1$	$0.3 \pm 0.1$
Temperate	pH control	$8.16 \pm 0.03$	$16.5 \pm 1.7$	$34 \pm 1$	$465 \pm 83$	$3.5 \pm 0.5$	$2.2 \pm 0.3$
	Moderate pH	$7.79 \pm 0.06$	$16.9 \pm 1.7$	$34 \pm 1$	$1130 \pm 12$	$1.6 \pm 0.0$	$1.0 \pm 0.0$
	Low pH	$7.62 \pm 0.05$	$16.6 \pm 1.7$	$34 \pm 1$	$1536 \pm 235$	$1.3 \pm 0.2$	$0.8 \pm 0.1$

Values for  $p\text{CO}_2$ ,  $\Omega$  calcite and  $\Omega$  aragonite were calculated from CO2SYS<sup>46</sup> with refitted constants<sup>47,48</sup>.

**Shell condition index.** Shell lengths were measured at the start and end of each experiment using Vernier calipers ( $\pm 0.1$  mm) to determine shell laid down in the natural environment that thickens from the internal surface as brachiopods grow (which we refer to as “thickening shell”) and shell growth extension during the experiments (which we refer to as “growing shell”). Mean lengths ( $\pm$  S.E.) of these two shell regions from each treatment are reported in Table S1. Scanning Electron Microscopes (JEOL 820 for the polar brachiopod and FEI QEMSCAN 650F for the temperate brachiopod; both operated using an accelerating voltage of 20 kV) were used to image gold-coated outer surfaces of five ventral valves of adult specimens from each treatment of both species to determine shell condition. Five types of shell

condition were present: intact shell (IS; intact periostracum with pitted layer), minimal wear (W1; periostracum without pitted layer), extensive wear (W2; wear but no dissolution), partial shell dissolution (SD1; dissolution in the inner primary layer) and extensive shell dissolution (SD2; dissolution exposing the innermost secondary layer). Full descriptions and examples of each type of shell condition for both species are presented in Table S2. Micrographs (1 mm x 1 mm) were collected at five standardised areas in thickening shell (areas located from umbo region towards anterior margin as detailed in Fig. S1A) and five standardised areas in growing shell (areas evenly spread in anterior margin as detailed in Fig. S1A). Percentage areas of each type of shell condition from each SEM micrograph were calculated/measured in ImageJ (Fig. S1B). Each shell region was analysed separately to determine whether treatment and/or the location of shell analysed (which we refer to as “shell position”) affected shell that had already been potentially subjected to substantial wear (thickening shell) and newly produced shell with less time subjected to wear (growing shell).

**Shell thickness.** Longitudinal cross sections of five dorsal valves of adult specimens from each treatment of both species were finely polished to 3  $\mu\text{m}$  using Kemet met papers (P400, P800, P2500 and P4000) followed by MetPrep diamond solutions (6  $\mu\text{m}$  and 3  $\mu\text{m}$ ). Acetate peels from polished cross sections of the brachial valves of both species were made according to a previous study<sup>53</sup>. Thickness measurements ( $\pm 0.1$  mm) of the primary layer, secondary layer and total shell were then measured from three areas of thickening shell (umbo region, middle of the shell

and nearer experimental growth as detailed in Fig. S2) and three areas of growing shell (oldest experimental growth to newest experimental growth in the anterior margin as detailed in Fig. S2) on a Swift monocular petrological microscope with fitted micrometer.

**Statistical analyses.** Shell condition index data were non-normally distributed due to the presence of zeros in the dataset. Non-parametric Kruskal-Wallis tests were, therefore, used to determine whether treatment and/or shell position affected the median percentage area of each type of shell condition. When significant differences occurred, post-hoc Dunn's tests were conducted to identify which treatments and shell positions were statistically different from each other. As shell condition and shell thickness measurements were conducted at several points within an individual, Kruskal-Wallis tests were also used to determine if individual number affected each shell condition. Linear mixed effects models were computed to determine if treatment, shell position (fixed effects) and/or individual number (random effect) impacted primary layer, secondary layer and total shell thickness:

$$\text{Thickness measurement} = \text{Treatment} + \text{Shell Position} + (1 | \text{Individual Number}) \\ + \text{error}$$

Likelihood ratio tests were used to determine p values ( $p < 0.05$ ) between the full model with the effect in question against the reduced model without the effect in question. When the ratio tests identified significant differences, post-hoc Tukey tests

228 were performed to determine which treatments or shell positions were responsible.  
229 Shell thickness data were checked for variance homogeneity and normality using  
230 Levene's and Shapiro-Wilk tests ( $p < 0.05$ ), respectively. Each shell region was  
231 analysed separately for both shell condition index and shell thickness to determine  
232 whether treatment and/or shell position affected shell maintenance (thickening shell)  
233 and shell production (growing shell). Statistical analyses were computed using R<sup>54</sup>  
234 with the *FSA* package<sup>55</sup> used for the Kruskal-Wallis and post-hoc Dunn's tests, the  
235 *lme4* package<sup>56</sup> for the linear mixed effects models and the *emmeans* package<sup>57</sup> for the  
236 post-hoc Tukey tests.

237

## RESULTS

**Shell condition index.**

*Thickening shell.* Intact shell (IS) was absent from both acidified treatments and only present in  $< 8.4 \pm 4.5\%$  (mean  $\pm$  SE) of both controls in the thickening shell in the polar brachiopod (Figure 1a & Figure 2). Instead, minimal wear (W1) dominated this region in both controls (Figure 2a;  $71.0 \pm 6.2\%$  in pH control and  $65.1 \pm 4.1\%$  in temperature control). Partial shell dissolution (SD1), however, was the most prominent shell condition in both acidified treatments (Figure 1a;  $64.3 \pm 4.6\%$  in moderate pH and  $71.7 \pm 4.3\%$  in low pH). With decreasing pH, the percentage area of partial shell dissolution increased (Figure 1a & Figure 2; Kruskal-Wallis:  $H = 70.93$ ,  $p < 0.001$ ). The extent of shell dissolution in the polar brachiopod also increased with decreasing pH (Kruskal-Wallis:  $H = 42.38$ ,  $p < 0.001$ ), with  $18.2 \pm 4.5\%$  of shell exhibiting exposed secondary layer in the low pH treatment (SD2) compared to  $0.9 \pm 0.4\%$  in the moderate pH treatment and the secondary layer never being exposed in either control. Temperature had no effect on shell dissolution or wear (Figure 1a; Dunn's Test: SD1 – T = 1.16,  $p = 0.244$ , SD2 – T = 0.26,  $p = 0.795$ , W1 – T = 0.25,  $p = 0.805$ , W2 – T = 1.60,  $p = 0.109$ ). In contrast to the polar brachiopod, thickening shell of the temperate brachiopod was mainly characterised by intact shell (Figure 1c; IS; 56.6 - 82.3%) across all treatments. Amounts of minimal wear (W1) decreased with decreasing pH in this shell region in the temperate brachiopod (Figure 1c; Kruskal-Wallis:  $H = 7.92$ ,  $p = 0.020$ ). Partial shell dissolution (SD1), however, increased with

260 decreasing pH (Figure 1c & Figure 3; Kruskal-Wallis:  $H = 53.72$ ,  $p < 0.001$ ) in growing  
261 shell in the temperate brachiopod. Shell dissolution in this temperate species was  
262 less extensive than for the polar species (Figure 2 & Figure 3) as the secondary layer  
263 was not exposed (SD2) in any individual in any treatment. Shell position or  
264 individual number did not affect any shell condition in the thickening shell of both  
265 species (Table S3).

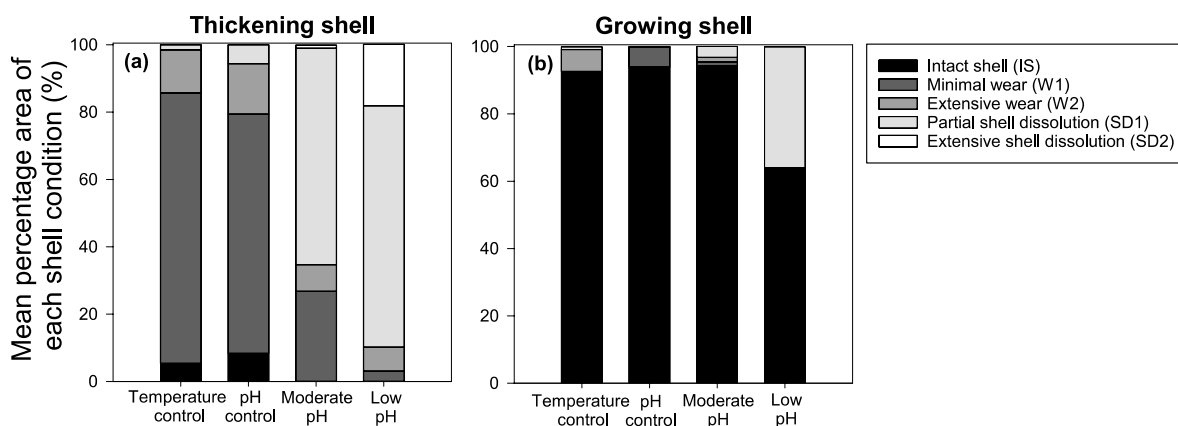
266 ***Growing shell.*** Growing shell in both species was mainly characterised by intact  
267 shell (IS) in all treatments (Figure 1b, d; polar brachiopod:  $> 63.9 \pm 4.7\%$ ; temperate  
268 brachiopod:  $> 83.2 \pm 1.8\%$ ). Less intact shell occurred in the most acidified conditions  
269 compared to all other treatments in both species (Figure 1b, d; Kruskal-Wallis: polar  
270 brachiopod -  $H = 41.81$ ,  $p < 0.001$ ; temperate brachiopod -  $H = 20.96$ ,  $p < 0.001$ ).  
271 Partial shell dissolution (SD1) increased with increasing acidity in the experimental  
272 growth of the polar brachiopod (Figure 1b & Figure 2; Kruskal-Wallis: polar  
273 brachiopod -  $H = 63.08$ ,  $p < 0.001$ ). This shell dissolution, however, occurred at a  
274 much lower level ( $3.2 \pm 1.0\%$  in moderate pH and  $35.9 \pm 4.7\%$  in low pH) in the  
275 growing shell than in the thickening shell in this species. Temperature had no effect  
276 on partial shell dissolution (Dunn's Test: Temperature control vs pH control:  $T = -$   
277  $0.22$ ,  $p = 0.829$ ). Partial shell dissolution (SD1) only occurred in the most acidified  
278 treatment in the temperate brachiopod (Figure 1d;  $11.1 \pm 1.5\%$ ), also in lower levels  
279 than in the thickening shell ( $28.3 \pm 3.2\%$ ) in this species. Extensive shell dissolution  
280 (SD2) was absent in the growing shell in both species across all treatments. Minimal  
281 wear (W1) was only present in two individuals across all treatments in the polar



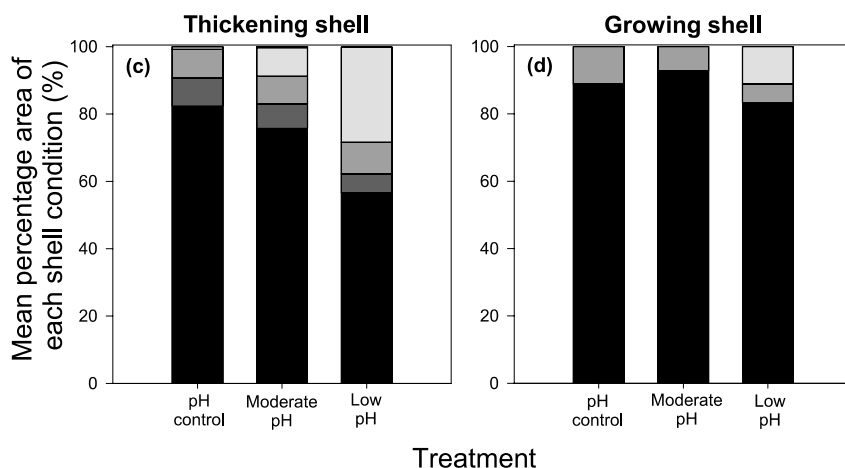
282 brachiopod and was absent from the temperate brachiopod. Extensive wear (W2)  
283 was present in higher levels in the control treatments of both species than in the  
284 acidified treatments (Figure 1c, d; Kruskal-Wallis; polar brachiopod –  $H = 43.98$ ,  $p <$   
285  $0.001$ , temperate brachiopod –  $H = 10.67$ ,  $p < 0.001$ ), however, only in low levels ( $<$   
286  $11.1 \pm 1.2\%$ ). Neither shell position nor individual number affected any shell  
287 condition in the growing shell of both species (Table S3).

288

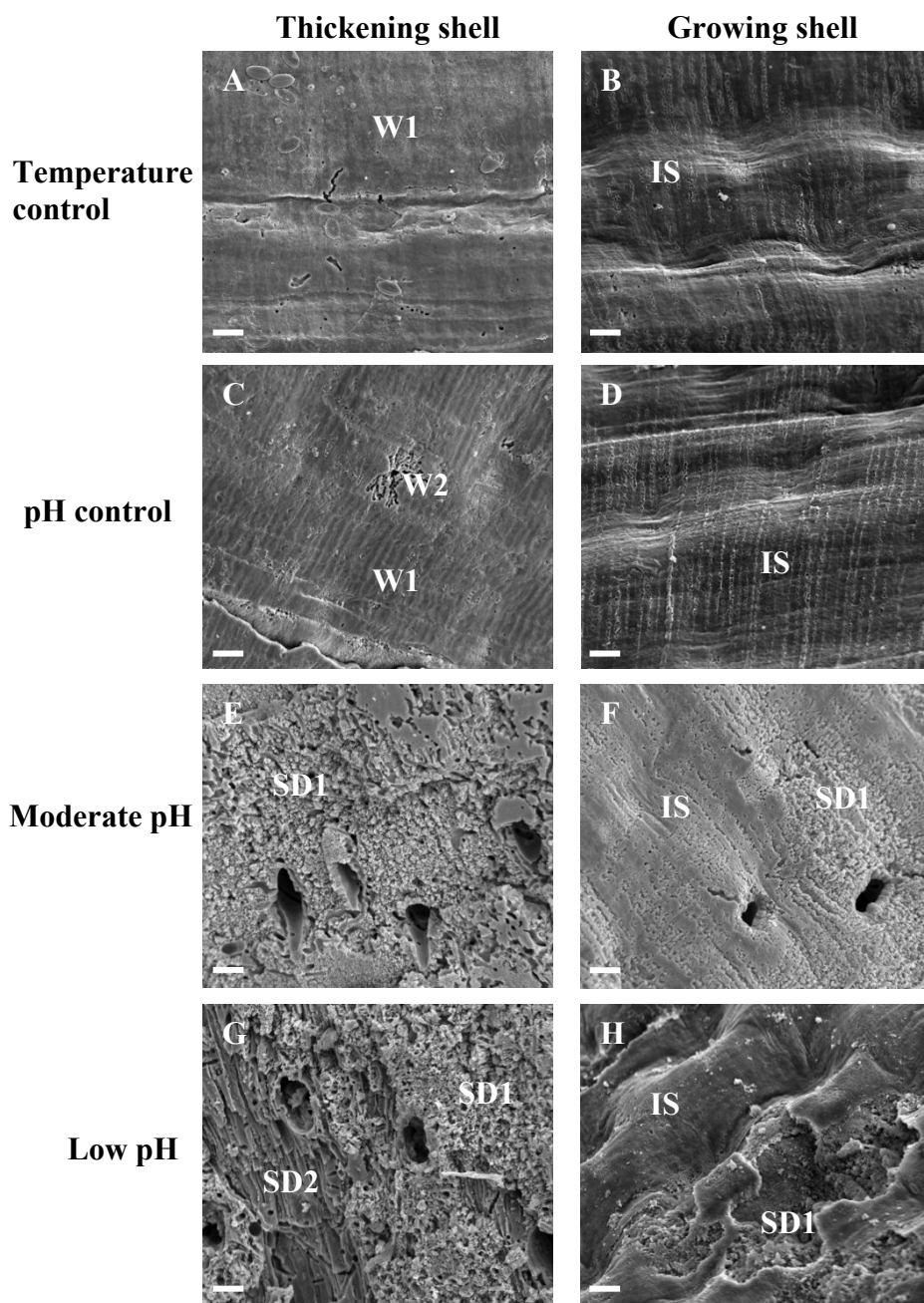
Polar brachiopod



Temperate brachiopod



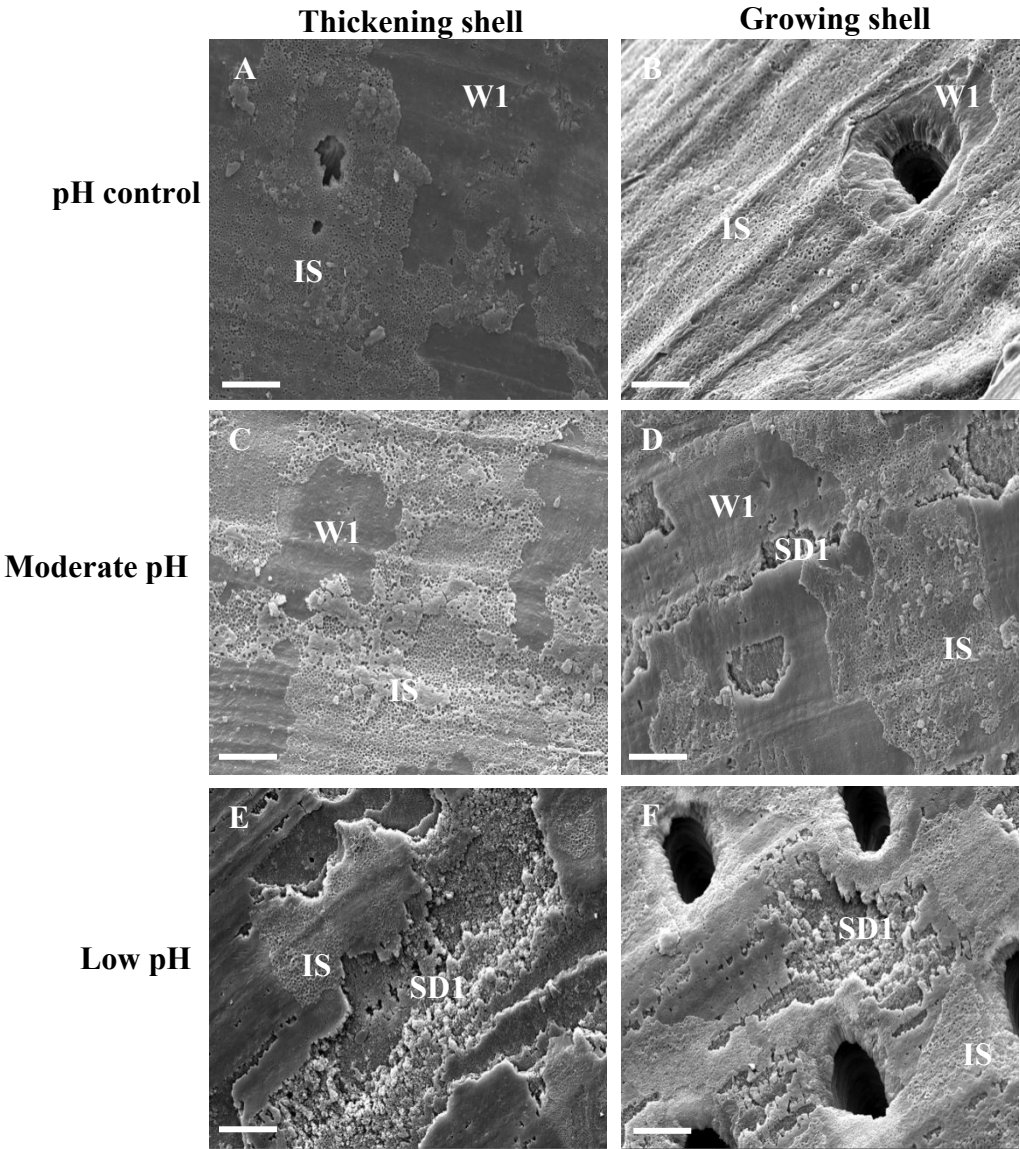
**Figure 1.** Representative shell condition - Mean percentage area of the different types of shell condition from five standardised areas in thickening shell (a, c) and five standardised areas in growing shell (b, d) in the polar brachiopod (top row, n = 5 per treatment) and in the temperate brachiopod (bottom row, n = 5 per treatment) in all treatments. Lighter grey tones indicate an increase in wear and/or shell dissolution (see legend).



296

297 **Figure 2.** Representative shell condition in the polar brachiopod – Examples of SEM  
 298 micrographs of shell surfaces of thickening shell (A, C, E, G) and growing shell (B, D,  
 299 F, H) in temperature control (A, B), pH control (C, D), moderate pH (E, F) and low  
 300 pH treatment (G, H). IS = intact shell, W1 = minimal wear, W2 = extensive wear, SD1  
 301 = partial shell dissolution and SD2 = extensive shell dissolution. Scale bar = 20  $\mu\text{m}$ .

302



303

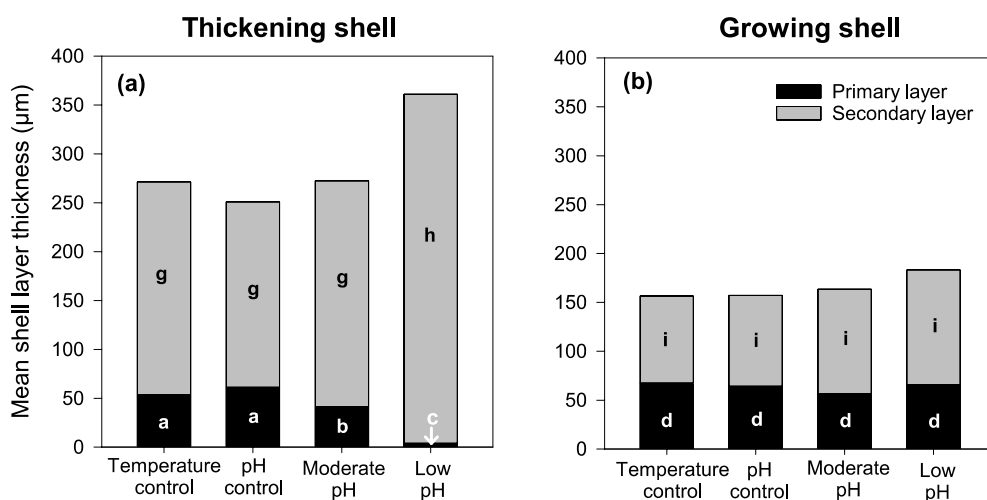
304 **Figure 3.** Representative shell condition in the temperate brachiopod – Examples of  
305 SEM micrographs of shell surfaces of thickening shell (A, C, E) and growing shell (B,  
306 D, F) in pH control (A, B), moderate pH (C, D) and low pH treatment (E, F). IS =  
307 intact shell, W1 = minimal wear, W2 = extensive wear and SD1 = partial shell  
308 dissolution. SD2 (extensive shell dissolution) was absent in all treatment in this  
309 species. Scale bar = 20  $\mu$ m.

**Shell thickness.**

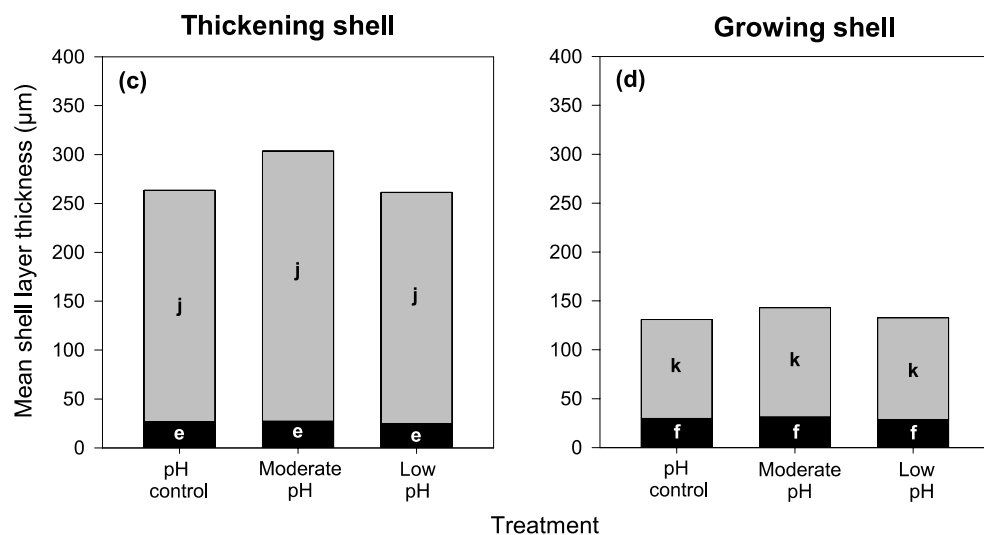
*Thickening shell.* The outer primary layer became progressively thinner in thickening shell as pH reduced in the polar brachiopod (Figure 4a; Linear Mixed Effects Model;  $\chi^2 = 79.72$ ,  $df = 3$ ,  $p < 0.001$ ). Secondary layer thickened in all treatments as this inner shell layer transitioned from the region of growing shell to thickening shell as the brachiopod grew (larger grey bars in Figure 4a vs smaller grey bars in Figure 4b). This inner secondary layer, and the whole shell, however, were thicker in the most acidified treatment in the thickening shell (Figure 4a; Linear Mixed Effects Model; Secondary Layer -  $\chi^2 = 39.63$ ,  $df = 3$ ,  $p < 0.001$ ; Total Shell -  $\chi^2 = 18.19$ ,  $df = 3$ ,  $p < 0.001$ ). Increased temperature had no effect on primary layer, secondary layer or total shell thickness (Tukey; Primary Layer -  $T = 1.73$ ,  $p = 0.319$ ; Secondary Layer -  $T = -1.20$ ,  $p = 0.627$ ; Total Shell -  $T = -0.80$ ,  $p = 0.855$ ). In contrast, neither individual shell layers nor total shell thickness were affected by lowered pH in the thickening shell in the temperate brachiopod (Figure 4c; Linear Mixed Effects Model; Primary Layer -  $\chi^2 = 4.17$ ,  $df = 2$ ,  $p = 0.124$ ; Secondary Layer -  $\chi^2 = 4.80$ ,  $df = 2$ ,  $p = 0.091$ ; Total Shell -  $\chi^2 = 4.27$ ,  $df = 2$ ,  $p = 0.118$ ). Primary layer was thinnest in the oldest part of the shell, the umbo region, across all treatments in both species (Table S4). Secondary layer and total shell thickness did not differ in different places in the thickening shell in each treatment in both species (Table S4). Individual number also had no effect on individual shell layer and total shell thickness in both species (Table S4).

332     *Growing shell.* Primary layer, secondary layer and total shell thickness were not  
333     impacted by lowered pH in either species (Figure 4b & 4d; Linear Mixed Effects  
334     Model; polar brachiopod: Primary Layer -  $\chi^2 = 3.62$ ,  $df = 3$ ,  $p = 0.306$ ; Secondary Layer  
335     -  $\chi^2 = 6.80$ ,  $df = 3$ ,  $p = 0.078$ ; Total Shell -  $\chi^2 = 5.26$ ,  $df = 3$ ,  $p = 0.154$ ; temperate  
336     brachiopod: Primary Layer -  $\chi^2 = 2.63$ ,  $df = 2$ ,  $p = 0.268$ ; Secondary Layer -  $\chi^2 = 0.82$ ,  $df$   
337     = 2,  $p = 0.663$ ; Total Shell -  $\chi^2 = 1.12$ ,  $df = 3$ ,  $p = 0.572$ ). Increased temperature also had  
338     no effect on either individual shell layers or total shell thickness in the polar  
339     brachiopod (Figure 4b). Primary layer thickness did not differ indifferent places  
340     throughout the growing shell in each treatment in either species (Table S4).  
341     Secondary layer and the total shell thickness, however, did get progressively thinner  
342     with the direction of growth in each treatment in both species (Table S4). Individual  
343     number also had no effect on individual shell layer and total shell thickness in  
344     growing shell in either species (Table S4).

## Polar brachiopod



## Temperate brachiopod



**Figure 4.** Shell thickness – Mean primary layer (black bar) and secondary layer (grey bar) thicknesses from three areas in the thickening shell (a, b) and from three areas in the growing shell (c, d) in the polar brachiopod (top row, n = 5 per treatment) and in the temperate brachiopod (bottom row, n = 5 per treatment) in all treatments. Whole bars represent total shell thickness. Lowercase letters a-f indicate significant differences in primary layer thickness and g-k represent significant differences in secondary layer and total shell thicknesses between treatments in each shell region

in each species. Comparisons were made only within shell region not between shell regions or between species.

## DISCUSSION

Long-term culturing of a polar and a temperate brachiopod under predicted end-century acidified conditions revealed that both species were more susceptible to shell dissolution with increasing acidity. Our two principal findings are significant dissolution and an unexpected compensation of induced thicker shells in the thickening shell.

**Dissolution of shell.** Shell loss has been widely reported in several marine calcifiers, however, these have largely been those which use higher solubility polymorphs of calcium carbonate (i.e. aragonite), such as corals<sup>58-60</sup> and molluscs<sup>23,32,61-64</sup>, high-magnesium calcite including coralline algae<sup>65,66</sup> and echinoderms<sup>67,68</sup>. Fewer studies have investigated shell dissolution in taxa which are entirely constructed of the lower solubility polymorph, low-magnesium calcite, such as rhynchonelliform brachiopods. Previously, the only other ocean acidification study assessing dissolution in brachiopods was conducted on dead shells<sup>38</sup>. Working on the polar species, they showed deterioration of the primary layer after only 35 days exposure to pH 7.4, which after 56 days exposed the fibres of the secondary layer below. This is the same dissolution pattern reported here in experiments involving live individuals after 7 months exposure to pH 7.54. Exposure of the



secondary layer calcite fibres may compromise shell integrity and probably strength due to the loss of the hard outer protective primary layer<sup>38,69</sup>.

Dissolution was more extensive in the polar than in the temperate brachiopod, as indicated by increased deterioration in the primary layer of the polar species compared to the temperate brachiopod in the moderate pH treatment and it was only in the polar species that the secondary layer was exposed in the low pH treatment. Antarctic calcified invertebrates are probably the most vulnerable organisms to ocean acidification for a number of reasons: they tend to be weakly calcified<sup>16,70</sup>; dissolution rates of calcium carbonate are inversely related to temperature<sup>17</sup>; and the polar regions are predicted to become the first to be undersaturated in aragonite by 2050 and calcite by 2100<sup>18,40,71-74</sup>. Both the moderate pH and the low pH treatment in the polar experiment were undersaturated with respect to calcite, however, both the acidified treatments in the temperate experiment were not undersaturated with respect to calcite. This could explain the differences in the extent of dissolution present between both investigated species. The state of the shells could have also influenced these species differences. Wear was more prominent in the thickening shell of the polar brachiopod than in the temperate brachiopod, which was most likely due to the longer lifespan of the polar species (up to 55-60 years)<sup>75</sup> compared to the temperate species (up to 14 years)<sup>76</sup>. Thus the shells of the polar brachiopod had been exposed to wear for a longer time in their natural environment before the experiment began. Such wear will have damaged or removed periostracum, which is key in protecting the animal from shell

396 dissolution<sup>31,77-82</sup>. Since periostracum is only formed at the growing edge of the  
397 mantle, it cannot be repaired if damaged or lost from the surface of the shell.  
398 Thinning or loss of this organic layer through physical or biotic abrasion and  
399 epibiont erosion, therefore, restricts protection from corrosive acidified waters. The  
400 periostracum in brachiopods is  $< 1 \mu\text{m}$  thick<sup>83</sup> and so is very vulnerable to loss.

401 Newly formed growing shell was mainly characterised by intact shell in both  
402 species. Partial shell dissolution did occur, however, in the most acidified treatment  
403 in both species albeit at a much lower level than in the thickening shell. Damage to  
404 the ultrathin periostracum from abrasion of other brachiopods in their conspecific  
405 cluster, natural decay of this outer layer or potentially the lowered pH conditions  
406 could have either softened the periostracum itself or disrupted the protective  
407 function of the periostracum. This latter possibility was suggested for external  
408 dissolution reported in newly formed shell in *M. edulis* after 2 months exposure to  
409  $1400 \mu\text{atm}$  and  $4000 \mu\text{atm}$ <sup>79</sup>. Disintegration of organic matrix in the shell rather than  
410 corrosion of crystals could have caused this shell degradation, as seen in spirorbids  
411 after 100-day exposure to pH 7.7 conditions<sup>84</sup>.

412 Temperature had no clear effect on shell dissolution or thickness in the polar  
413 brachiopod as indicated by the lack of or only minimal primary layer dissolution  
414 and no change in any thickness measurement in both thickening and growing shell  
415 in the temperature control (held at  $0^{\circ}\text{C}$  – current average Antarctic summer  
416 temperatures) and the pH control (kept at the  $2^{\circ}\text{C}$  temperature increase predicted for  
417 2100). In contrast, temperature and not acidification reduced shell strength in *M.*

*edulis* after 6 months exposure to forecasted end-century pH and warming conditions<sup>85</sup>. It was concluded that warming had an indirect effect on shell strength by shifting the energy budget from shell deposition to increased maintenance costs. Food availability was limited throughout the experiment, which would likely have enhanced the temperature effect as low food levels can reduce shell growth and significantly influence the amount of inner shell dissolution in *M. edulis* after 7 weeks exposure to varying  $p\text{CO}_2$  levels<sup>86</sup>. This highlights the necessity of using multistressors in ocean acidification research to better understand the abilities of marine calcifiers to maintain shell integrity under future predicted environmental conditions.

**Compensation.** Despite the widely reported significant effects of dissolution on marine calcifiers in ocean acidification research, very few studies investigate organisms' abilities to compensate for shell loss. New shell deposited by *M. edulis* after 9 months exposure to 750  $\mu\text{atm}$  and 1000  $\mu\text{atm}$   $p\text{CO}_2$  was rounder and flatter with a thinner aragonite layer than shell produced in ambient conditions of 380  $\mu\text{atm}$ <sup>27</sup>. The authors attributed this new shell shape to a compensatory mechanism to enhance protection from predators and changing environments as these mussels were unable to grow thicker shells in high  $p\text{CO}_2$  conditions. Shell thickening has occurred in response to biotic shell loss by endoliths and other conspecifics grazing on their external shell in Patellid limpets, *Patella granatina* and *P. argenvillei*<sup>87</sup>, and to abiotic shell loss by physical impacts from ice in the Antarctic limpet *Nacella concinna*<sup>88</sup>. Decreased shell thickness has also been reported in molluscs in lowered

pH conditions, due to internal dissolution of the highly soluble aragonite layer<sup>27,78,86</sup>. For compensatory mechanisms to succeed, they must occur at faster rates than the deleterious effect. Thicker basal shells were reported in the barnacle *Amphibalanus amphitrite* under lowered pH conditions (pH 7.4), however, this compensation calcification was insufficient as dissolution weakened shells faster than it was deposited<sup>89</sup>. A pteropod specimen collected from the Fram Strait in the Arctic Ocean also produced a shell four times thicker than the original shell in response to mechanical and dissolution damage from undersaturated waters<sup>31</sup>.

Extensive shell dissolution at low pH in thickening shell of the polar brachiopod led to a drastic decrease in primary layer thickness. The polar species counteracted this chemical attack by laying down more secondary layer on the internal surface of the shell, which resulted in increased overall shell thickness during the experimental period. The less extensive dissolution in the temperate brachiopod was reflected by no clear impact of acidified conditions on either total shell or individual shell layer thicknesses. Our findings appear to contrast with reports of primary layer thickening in the Chilean terebratulide *Magellania venosa* after being cultured in pH 7.35 conditions<sup>90</sup>, however, their observations appear to be based on only one measured specimen in both the acidified treatment and the control.

Compensatory mechanisms must also be sufficient in maintaining an organism's overall performance. The secondary layer of terebratulide brachiopod shells is softer than the harder protective primary layer<sup>37,91</sup> raising the question of whether a shell made solely out of secondary layer would provide adequate protection to ensure

survival. No external dissolution of the exposed secondary layer of the polar brachiopod was observed perhaps due to protection from the organic matrix shrouding calcite crystals of this innermost fibrous shell layer<sup>92,93</sup>. Primary layer is often missing in older parts of brachiopod shells or in older individuals<sup>90</sup>, therefore, a thicker shell consisting of only secondary layer could provide sufficient protection in predicted pH conditions expected by 2100. Although, ocean acidification impacts on brachiopod shell strength warrant further investigation.

Total shell thickness or individual shell layer thickness of growing shell of both species were not affected by predicted end-century acidified conditions. Shell thickness, therefore, is only impacted by lowered pH when extensive shell dissolution occurs. In a previous study, shell thickness in the temperate brachiopod did not vary over the last 120 years despite a 0.1 pH unit decrease and 2°C increase in temperature since the Industrial Revolution<sup>94</sup>. Forecasted acidified conditions by 2100 also did not impact shell growth rates and the ability to shell repair in both the polar and temperate brachiopod<sup>6,7</sup>. The resilience of shell thickness in both the polar and temperate species to past and predicted environmental change, in addition to their unaffected shell growth rates under end-century pH levels<sup>6,7</sup>, indicates the robust ability of rhynchonelliform brachiopods to construct shell under acidified conditions. The thickness of calcite and aragonite layers in newly formed shell of *M. edulis* were also not affected by elevated  $p\text{CO}_2$ <sup>79</sup>. This lack of variation in shell thickness to acidified conditions in newly produced shell further demonstrates the increase of shell thickness in the thickening shell is a compensatory response to

extensive shell dissolution occurring at the external shell surface, although the mechanisms whereby the brachiopods identify the shell is thinning remain to be elucidated.

The extent of vulnerability of two highly calcium-carbonate-dependent species to dissolution in acidified seawater is concerning. Without any counteracting response, dissolution may compromise shell integrity leading to reduced protection and decreased suitability of brachiopod shells as a habitat for other marine organisms. Physiological acclimatisation is one approach organisms can utilise to cope with such threats in the challenging conditions predicted by 2100. We identified induced shell thickening forming thicker shells in the polar brachiopod as a compensatory mechanism to extensive shell dissolution under lowered pH levels. The less extensive dissolution in the temperate species was probably a function of higher temperatures in the temperate study and the corresponding lower  $\text{CaCO}_3$  solubility. This suggests that the level of dissolution in the temperate brachiopod after 3 months exposure to predicted end-century pH conditions did not induce similar compensation. This induced shell thickening could come at an overall cost to the organism as increased shell production is energy-demanding, involving the accumulation, transportation and precipitation of calcium carbonate as well as the production of the organic matrix<sup>95,96</sup>. Acidification also significantly increases the proportion of the animal's energy budget that needs to be devoted to shell production<sup>97</sup>, therefore, there may be long-term impacts on life histories and maintenance of populations. Long-term experiments investigating the capacity of

organisms to acclimatise and possibly adapt to future change is crucial to further our understanding of how marine organisms will cope with future climate change.

Marine organisms may also adjust physiological, behavioural or ecological traits as additional compensatory responses to their changing habitats. As well as direct effects on energy budgets (e.g. induced shell thickening), ocean acidification could also have indirect impacts through the alteration of their resource quality (e.g. energy intake)<sup>98</sup>. To maintain organismal homeostasis in varying environments, individuals may compensate by modifying the quality and quantity of food consumed, which in turn could also stabilise community productivity<sup>99</sup>. Multiple compensatory mechanisms could be paramount to maintain overall performance of organisms and subsequently sustain key community processes under future environmental change.

## ASSOCIATED CONTENT

### **Supporting Information**

Mean lengths ( $\pm$ S.E) of thickening and growing shell regions (Table S1)

Descriptions and examples of each shell condition (Table S2)

Schematic and example of shell condition index measurements (Figure S1)

Schematic and example of shell thickness measurements (Figure S2)

Shell dissolution additional statistical results (Table S3)

Shell thickness additional statistical results (Table S4)

527

528 AUTHOR INFORMATION

529 **Corresponding Author**

530 \* Emma L. Cross. Email: E.L.Cross@cantab.net. Address: Department of Earth  
531 Sciences, University of Cambridge, Downing Street, Cambridge, CB2 3EQ, UK and  
532 British Antarctic Survey, Natural Environment Research Council, High Cross,  
533 Madingley Road, Cambridge, CB3 0ET, UK

534 ACKNOWLEDGEMENTS

535 The authors would like to thank the scientific dive team at the British Antarctic  
536 Survey Rothera Research Station for collecting the *L. uva* specimens and to Dr.  
537 Coleen Suckling, Dr. Simon Morley and Rebecca Smith for their help in the set-up  
538 and maintenance of the polar experimental system. We are very grateful to Dr. Miles  
539 Lamare for his assistance in the organisation and collection of the *C. inconspicua*  
540 specimens. Thanks also to the science support staff at the Portobello Marine  
541 Laboratory, University of Otago for their help in the construction and maintenance  
542 of the temperate experimental system. Thanks to Dr. Kim Currie at the National  
543 Institute of Water and Atmospheric Research for the DIC and total alkalinity  
544 measurements. This research was funded by the NERC PhD Studentship  
545 (NE/T/A/2011) awarded to ELC.

546

547



## REFERENCES

- (1) Peck, L. S., Organisms and responses to environmental change. *Mar. Genom.* **2011**, *4*, 237-243.
- (2) Doney, S. C.; Fabry, V. J.; Feely, R. A.; Kleypas, J. A., Ocean acidification: the other CO<sub>2</sub> problem. *Ann. Rev. Mar. Sci.* **2009**, *1*, 169-192.
- (3) Hofmann, G. E.; Barry, J. P.; Edmunds, P. J.; Gates, R. D.; Hutchins, D. A.; Klinger, T.; Sewell, M. A., The effect of ocean acidification on calcifying organisms in marine ecosystems: An organism-to-ecosystem perspective. *Annu Rev Ecol Evol Syst* **2010**, *41*, 127-147.
- (4) Byrne, M., Impact of ocean warming and ocean acidification on marine invertebrate life history stages: vulnerabilities and potential for persistence in a changing ocean. *Oceanogr Mar Biol Annu Rev* **2011**, *49*, 1-42.
- (5) Parker, L. M.; Ross, P. M.; O'Connor, W. A.; Pörtner, H. O.; Scanes, E.; Wright, J. M., Predicting the response of molluscs to the impact of ocean acidification. *Biology* **2013**, *2*, 651-92.
- (6) Cross, E. L.; Peck, L. S.; Harper, E. M., Ocean acidification does not impact shell growth or repair of the Antarctic brachiopod *Liothyrella uva* (Broderip, 1833). *J. Exp. Mar. Biol. Ecol.* **2015**, *462*, 29-35.
- (7) Cross, E. L.; Peck, L. S.; Lamare, M. D.; Harper, E. M., No ocean acidification effects on shell growth and repair in the New Zealand brachiopod *Calloria inconspicua* (Sowerby, 1846). *ICES J. Mar. Sci.* **2016**, *73*, 920-926.
- (8) Queirós, A. M.; Fernandes, J. A.; Faulwetter, S.; Nunes, J.; Rastrick, S. P.; Mieszkowska, N.; Artioli, Y.; Yool, A.; Calosi, P.; Arvanitidis, C.; Findlay, H. S.; Barange, M.; Cheung, W. W.; Widdicombe, S., Scaling up experimental ocean acidification and warming research: from individuals to the ecosystem. *Global Change Biol.* **2015**, *21*, 130-143.
- (9) Suckling, C. C.; Clark, M. S.; Richard, J.; Morley, S. A.; Thorne, M. A.; Harper, E. M.; Peck, L. S., Adult acclimation to combined temperature and pH

576 stressors significantly enhances reproductive outcomes compared to short-term  
577 exposures. *J. Anim. Ecol.* **2014**, *84*, 773-784.

578 (10) Hazan, Y.; Wangensteen, O. S.; Fine, M., Tough as a rock-boring urchin:  
579 adult *Echinometra* sp. EE from the Red Sea show high resistance to ocean  
580 acidification over long-term exposures. *Mar. Biol.* **2014**, *161*, 2531-2545.

581 (11) Collins, S.; Rost, B.; Rynearson, T. A., Evolutionary potential of marine  
582 phytoplankton under ocean acidification. *Evol. Appl.* **2014**, *7*, 140-155.

583 (12) Kelly, M. W.; Hofmann, G. E., Adaptation and the physiology of ocean  
584 acidification. *Funct. Ecol.* **2013**, *27*, 980-990.

585 (13) Sunday, J. M.; Calosi, P.; Dupont, S.; Munday, P. L.; Stillman, J. H.;  
586 Reusch, T. B., Evolution in an acidifying ocean. *Trends Ecol. Evol.* **2014**, *29*, 117-125.

587 (14) Leung, J. Y. S.; Russell, B. D.; Connell, S. D., Mineralogical plasticity acts  
588 as a compensatory mechanism to the impacts of ocean acidification. *Environ. Sci.*  
589 *Technol.* **2017**, *51*, 2652-2659.

590 (15) Byrne, M.; Przeslawski, R., Multistressor impacts of warming and  
591 acidification of the ocean on marine invertebrates' life histories. *Integ. Comp. Biol.*  
592 **2013**, *53*, 582-596.

593 (16) Watson, S.-A.; Peck, L. S.; Tyler, P. A.; Southgate, P. C.; Tan, K. S.; Day, R.  
594 W.; Morley, S. A., Marine invertebrate skeleton size varies with latitude, temperature  
595 and carbonate saturation: implications for global change and ocean acidification.  
596 *Global Change Biol.* **2012**, *18*, 3026-3038.

597 (17) Revelle, R. R.; Fairbridge, R. W., Carbonates and carbon dioxide. In  
598 *Geological Society of America Memoirs*, 1957; Vol. 67, pp 239-296.

599 (18) Fabry, V. J.; McClintock, J. B.; Mathis, J. T.; Grebmeier, J. M., Ocean  
600 acidification at high latitudes: the bellwether. *Oceanography* **2009**, *22*, 160-171.

601 (19) Vermeij, G. J., The Mesozoic marine revolution: evidence from snails,  
602 predators and grazers. *Paleobiology* **1977**, *3*, 245-258.

(20) Harper, E. M.; Clark, M. S.; Hoffman, J. I.; Philipp, E. E.; Peck, L. S.; Morley, S. A., Iceberg scour and shell damage in the Antarctic bivalve *Laternula elliptica*. *PLoS ONE* **2012**, *7*, e46341.

(21) Harper, E. M., The molluscan periostracum: an important constraint in bivalve evolution. *Palaeontology* **1997**, *40*, 71-97.

(22) Williams, A.; Mackay, S., Secretion and ultrastructure of the periostracum of some terebratulide brachiopods. *Proc. R. Soc. B.* **1978**, *202*, 191-209.

(23) Nienhuis, S.; Palmer, A. R.; Harley, C. D., Elevated CO<sub>2</sub> affects shell dissolution rate but not calcification rate in a marine snail. *Proc. R. Soc. B.* **2010**, *277*, 2553-2558.

(24) Bausch, A. R.; Gallego, M. A.; Harianto, J.; Thibodeau, P.; Bednaršek, N.; Havenhand, J. N.; Klinger, T., Influence of bacteria on shell dissolution in dead gastropod larvae and adult *Limacina helicina* pteropods under ocean acidification conditions. *Mar. Biol.* **2018**, *165*, 40.

(25) Freeman, A. S.; Byers, J. E., Divergent induced responses to an invasive predator in marine mussel populations. *Science* **2006**, *313*, 831-834.

(26) Peyer, S. M.; Hermanson, J. C.; Lee, C. E., Developmental plasticity of shell morphology of quagga mussels from shallow and deep-water habitats of the Great Lakes. *J. Exp. Biol.* **2010**, *213*, 2602-2609.

(27) Fitzer, S. C.; Vittert, L.; Bowman, A.; Kamenos, N. A.; Phoenix, V. R.; Cusack, M., Ocean acidification and temperature increase impact mussel shell shape and thickness: problematic for protection? *Ecol. Evol.* **2015**, *5*, 4875-4884.

(28) Telesca, L.; Michalek, K.; Sanders, T.; Peck, L. S.; Thyrring, J.; Harper, E. M., Blue mussel shell shape plasticity and natural environments: a quantitative approach. *Sci. Rep.* **2018**, *8*, 2865.

(29) Vermeij, G. J., Phenotypic evolution in a poorly dispersing snail after arrival of a predator. *Nature* **1982**, *299*, 349-350.

(30) Fisher, J. A.; Rhile, E. C.; Liu, H.; Petraitis, P. S., An intertidal snail shows a dramatic size increase over the past century. *PNAS* **2009**, *106*, 5209-12.

- 632 (31) Peck, V. L.; Oakes, R. L.; Harper, E. M.; Manno, C.; Tarling, G. A.,  
633 Pteropods counter mechanical damage and dissolution through extensive shell  
634 repair. *Nat. Commun.* **2018**, *9*, 264.
- 635 (32) Duquette, A.; McClintock, J. B.; Amsler, C. D.; Pérez-Huerta, A.; Milazzo,  
636 M.; Hall-Spencer, J. M., Effects of ocean acidification on the shells of four  
637 Mediterranean gastropod species near a CO<sub>2</sub> seep. *Mar. Pollut. Bull.* **2017**.
- 638 (33) Telesca, L.; Peck, L. S.; Sanders, T.; Thyrring, J.; Sejr, M. K.; Harper, E. M.,  
639 Plasticity and environmental heterogeneity predict geographic resilience patterns of  
640 foundation species to future change. *bioRxiv* **2018**.
- 641 (34) Peck, L. S., The tissues of articulate brachiopods and their value to  
642 predators. *Phil. Trans. R. Soc. B.* **1993**, *339*, 17-32.
- 643 (35) Peck, L. S., Brachiopods and climate change. *Earth Env. Sci. T. R. So.* **2008**,  
644 *98*, 451-456.
- 645 (36) Williams, A.; Brunton, C. H. C.; MacKinnon, D. I., Morphology. In *Treatise*  
646 *on Invertebrate Paleontology, Part H, Brachiopods (Revised)*, Kaesler, R. L., Ed. The  
647 Geological Society of America and The University of Kansas Press: Boulder,  
648 Colorado, and Lawrence, Kansas, 1997; Vol. 1, pp 321-422.
- 649 (37) Goetz, A. J.; Griesshaber, E.; Neuser, R. D.; Lüter, C.; Hühner, M.; Harper,  
650 E. M.; Schmahl, W. W., Calcite morphology, texture and hardness in the distinct  
651 layers of rhynchonelliform brachiopod shells. *Eur. J. Mineral.* **2009**, *21*, 303-315.
- 652 (38) McClintock, J. B.; Angus, R. A.; McDonald, M. R.; Amsler, C. D.;  
653 Catledge, S. A.; Vohra, Y. K., Rapid dissolution of shells of weakly calcified Antarctic  
654 benthic macroorganisms indicates high vulnerability to ocean acidification. *Antarct.*  
655 *Sci.* **2009**, *21*, 449-456.
- 656 (39) Venables, H. J.; Clarke, A.; Meredith, M. P., Wintertime controls on  
657 summer stratification and productivity at the western Antarctic Peninsula. *Limnol.*  
658 *Oceanogr.* **2013**, *58*, 1035-1047.
- 659 (40) McNeil, B. I.; Matear, R. J., Southern Ocean acidification: a tipping point  
660 at 450-ppm atmospheric CO<sub>2</sub>. *PNAS* **2008**, *105*, 18860-18864.

(41) Greig, M. J.; Ridgway, N. M.; Shakespeare, B. S., Sea surface temperature variations at coastal sites around New Zealand. *N. Z. J. Mar. Freshwat. Res.* **1988**, *22*, 391-400.

(42) Roper, D. S.; Jillett, J. B., Seasonal occurrence and distribution of flatfish (Pisces: Pleuronectiformes) in inlets and shallow water along the Otago coast. *N. Z. J. Mar. Freshwat. Res.* **1981**, *15*, 1-13.

(43) IPCC, Climate Change 2013: The Physical Science Basis. In *Working Group I Contribution to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, Stocker, T. F.; Qin, D.; Plattner, G.-K.; Tignor, M.; Allen, S. K.; Boschung, J.; Nauels, A.; Xia, Y.; Bex, V.; Midgley, P. M., Eds. Cambridge, United Kingdom and New York, NY, USA, 2013; p 1552.

(44) Mitchell, J. F. B.; Senior, C. A.; Johns, T. C., *Transient response to increasing greenhouse gases using models with and without flux adjustment*. Metrological Office: 1998.

(45) Clarke, A.; Meredith, M. P.; Wallace, M. I.; Brandon, M. A.; Thomas, D. N., Seasonal and interannual variability in temperature, chlorophyll and macronutrients in northern Marguerite Bay, Antarctica. *Deep Sea Res. Part II Top. Stud. Oceanogr.* **2008**, *55*, 1988-2006.

(46) Lewis, E.; Wallace, D. W. R.; Allison, L. J., *Program developed for CO<sub>2</sub> system calculations*. Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory, US Department of Energy, Oak Ridge, Tennessee, US, 1998.

(47) Mehrbach, C.; Culberson, C. H.; Hawley, J. E.; Pytkowicz, R. M., Measurement of apparent dissociation constants of carbonic acid in seawater at atmospheric pressure. *Limnol. Oceanogr.* **1973**, *18*, 897-907.

(48) Dickson, A. G.; Millero, F. J., A comparison of the equilibrium-constants for the dissociation of carbonic-acid in seawater media. *Deep-Sea Res Pt I* **1987**, *34*, 1733-1743.

- 688 (49) Garibotti, I. A.; Vernet, M.; Ferrario, M. E., Annually recurrent  
689 phytoplanktonic assemblages during summer in the seasonal ice zone west of the  
690 Antarctic Peninsula (Southern Ocean). *Deep-Sea Res Pt I* **2005**, *52*, 1823-1841.
- 691 (50) Garibotti, I. A.; Vernet, M.; Ferrario, M. a. E.; Smith, R. C.; Ross, R. M.;  
692 Quentin, L. B., Phytoplankton spatial distribution patterns along the western  
693 Antarctic Peninsula (Southern Ocean). *Mar. Ecol. Prog. Ser.* **2003**, *261*, 21-39.
- 694 (51) Dickson, A. G.; Sabine, C. L.; Christian, J. R. *Guide to best practices for*  
695 *Ocean CO<sub>2</sub> Measurements*; 2007.
- 696 (52) Wanninkhof, R.; Lewis, E.; Feely, R. A.; Millero, F. J., The optimal  
697 carbonate dissociation constants for determining surface water  $p\text{CO}_2$  from alkalinity  
698 and total inorganic carbon. *Mar. Chem.* **1999**, *65*, 291-301.
- 699 (53) Richardson, C. A.; Crisp, D. J.; Runham, N. W., Tidally deposited growth  
700 bands in the shell of the common cockle, *Cerastoderma edula* (L.). *Malacologia* **1979**, *18*,  
701 277-290.
- 702 (54) R, Core Team, R: A language and environment for statistical computing.  
703 R Foundation for Statistical Computing, Vienna, Austria. URL [https://www.R-](https://www.R-project.org/)  
704 [project.org/](https://www.R-project.org/). **2017**.
- 705 (55) Ogle, D. H., Fisheries Stock Analysis. **2018**, R Package version 0.8.20.
- 706 (56) Bates, D.; Maechler, M.; Bolker, B.; Walker, S., Fitting linear mixed-effects  
707 models using lme4. *J. Stat. Softw.* **2015**, *67*, 1-48.
- 708 (57) Lenth, R., emmeans: Estimated Marginal Means, aka Least-Squares  
709 Means. **2018**, R package version 1.2.3.
- 710 (58) Comeau, S.; Carpenter, R. C.; Lantz, C. A.; Edmunds, P. J., Ocean  
711 acidification accelerates dissolution of experimental coral reef communities.  
712 *Biogeosciences* **2015**, *12*, 365-372.
- 713 (59) Silbiger, N. J.; Donahue, M. J., Secondary calcification and dissolution  
714 respond differently to future ocean conditions. *Biogeosciences* **2015**, *12*, 567-578.
- 715 (60) Andersson, A. J.; Kuffner, I. B.; MacKenzie, F. T.; Jokiel, P. L.; Rodgers, K.  
716 S.; Tan, A., Net loss of  $\text{CaCO}_3$  from a subtropical calcifying community due to

717 seawater acidification: mesocosm-scale experimental evidence. *Biogeosciences* **2009**, *6*,  
718 1811-1823.

719 (61) Hall-Spencer, J. M.; Rodolfo-Metalpa, R.; Martin, S.; Ransome, E.; Fine,  
720 M.; Turner, S. M.; Rowley, S. J.; Tedesco, D.; Buia, M. C., Volcanic carbon dioxide  
721 vents show ecosystem effects of ocean acidification. *Nature* **2008**, *454*, 96-9.

722 (62) Milano, S.; Schöne, B. R.; Wang, S.; Müller, W. E., Impact of high  $p\text{CO}_2$  on  
723 shell structure of the bivalve *Cerastoderma edule*. *Mar. Environ. Res.* **2016**, *119*, 144-155.

724 (63) Bednaršek, N.; Tarling, G. A.; Bakker, D. C. E.; Fielding, S.; Feely, R. A.,  
725 Dissolution dominating calcification process in polar pteropods close to the point of  
726 aragonite undersaturation. *PLoS ONE* **2014**, *9*, e109183.

727 (64) Harvey, B. P.; Agostini, S.; Wada, S.; Inaba, K.; Hall-Spencer, J. M.,  
728 Dissolution: The achilles' heel of the triton shell in an acidifying ocean. *Front. Mar.*  
729 *Sci.* **2018**, *5*, 371.

730 (65) Kamenos, N. A.; Burdett, H. L.; Aloisio, E.; Findlay, H. S.; Martin, S.;  
731 Longbone, C.; Dunn, J.; Widdicombe, S.; Calosi, P., Coralline algal structure is more  
732 sensitive to rate, rather than the magnitude, of ocean acidification. *Global Change Biol.*  
733 **2013**, *19*, 3621-3628.

734 (66) Cornwall, C. E.; Boyd, P. W.; McGraw, C. M.; Hepburn, C. D.; Pilditch, C.  
735 A.; Morris, J. N.; Smith, A. M.; Hurd, C. L., Diffusion boundary layers ameliorate the  
736 negative effects of ocean acidification on the temperate coralline macroalga  
737 *Arthrocardia corymbosa*. *PLoS ONE* **2014**, *9*, e97235.

738 (67) Miles, H.; Widdicombe, S.; Spicer, J. I.; Hall-Spencer, J., Effects of  
739 anthropogenic seawater acidification on acid-base balance in the sea urchin  
740 *Psammechinus miliaris*. *Mar. Pollut. Bull.* **2007**, *54*, 89-96.

741 (68) Dubois, P., The skeleton of postmetamorphic echinoderms in a changing  
742 world. *Biol. Bull.* **2014**, *226*, 223-236.

743 (69) Scurr, D. J.; Eichhorn, S. J., Structure/property relationships in seashells.  
744 In *Mechanical Properties of Bioinspired Materials* Viney, C.; Katti, K.; Ulm, F.-J.;

- 745 Hellmich, C., Eds. Materials Research Society Symposium Proceedings: Warrendale,  
746 PA, 2005; pp 87–92.
- 747 (70) Nicol, D., Some characteristics of cold-water marine pelecypods. *J.*  
748 *Paleontol.* **1967**, *41*, 1330-1340.
- 749 (71) Caldeira, K.; Wickett, M. E., Ocean model predictions of chemistry  
750 changes from carbon dioxide emissions to the atmosphere and ocean. *J. Geophys. Res.*  
751 **2005**, *110*, C09S04.
- 752 (72) Orr, J. C.; Fabry, V. J.; Aumont, O.; Bopp, L.; Doney, S. C.; Feely, R. A.;  
753 Gnanadesikan, A.; Gruber, N.; Ishida, A.; Joos, F.; Key, R. M.; Lindsay, K.; Maier-  
754 Reimer, E.; Matear, R.; Monfray, P.; Mouchet, A.; Najjar, R. G.; Plattner, G. K.;  
755 Rodgers, K. B.; Sabine, C. L.; Sarmiento, J. L.; Schlitzer, R.; Slater, R. D.; Totterdell, I.  
756 J.; Weirig, M. F.; Yamanaka, Y.; Yool, A., Anthropogenic ocean acidification over the  
757 twenty-first century and its impact on calcifying organisms. *Nature* **2005**, *437*, 681-6.
- 758 (73) Fabry, V. J.; Seibel, B. A.; Feely, R. A.; Orr, J. C., Impacts of ocean  
759 acidification on marine fauna and ecosystem processes. *ICES J. Mar. Sci.* **2008**, *65*,  
760 414-432.
- 761 (74) Guinotte, J. M.; Fabry, V. J., Ocean acidification and its potential effects on  
762 marine ecosystems. *Ann. N. Y. Acad. Sci.* **2008**, *1134*, 320-42.
- 763 (75) Peck, L. S.; Brey, T., Bomb signals in old Antarctic brachiopods. *Nature*  
764 **1996**, *380*, 207-208.
- 765 (76) Doherty, P. J., A demographic study of a subtidal population of the New  
766 Zealand articulate brachiopod *Terebratella inconspicua*. *Mar. Biol.* **1979**, *52*, 331-342.
- 767 (77) Ries, J. B.; Cohen, A. L.; McCorkle, D. C., Marine calcifiers exhibit mixed  
768 responses to CO<sub>2</sub>-induced ocean acidification. *Geology* **2009**, *37*, 1131-1134.
- 769 (78) Tunnicliffe, V.; Davies, K. T. A.; Butterfield, D. A.; Embley, R. W.; Rose, J.  
770 M.; Chadwick Jr, W. W., Survival of mussels in extremely acidic waters on a  
771 submarine volcano. *Nat. Geosci.* **2009**, *2*, 344-348.
- 772 (79) Thomsen, J.; Gutowska, M. A.; Saphörster, J.; Heinemann, A.;  
773 Trübenbach, K.; Fietzke, J.; Hiebenthal, C.; Eisenhauer, A.; Körtzinger, A.; Wahl, M.;



Melzner, F., Calcifying invertebrates succeed in a naturally CO<sub>2</sub>-rich coastal habitat but are threatened by high levels of future acidification. *Biogeosciences* **2010**, *7*, 3879-3891.

(80) Rodolfo-Metalpa, R.; Houlbreque, F.; Tambutte, E.; Boisson, F.; Baggini, C.; Patti, F. P.; Jeffree, R.; Fine, M.; Foggo, A.; Gattuso, J.-P.; Hall-Spencer, J. M., Coral and mollusc resistance to ocean acidification adversely affected by warming. *Nature Clim. Change* **2011**, *1*, 308-312.

(81) Coleman, D. W.; Byrne, M.; Davis, A. R., Molluscs on acid: gastropod shell repair and strength in acidifying oceans. *Mar. Ecol. Prog. Ser.* **2014**, *509*, 203-211.

(82) Peck, V. L.; Tarling, G. A.; Manno, C.; Harper, E. M.; Tynan, E., Outer organic layer and internal repair mechanism protects pteropod *Limacina helicina* from ocean acidification. *Deep Sea Res. Part II Top. Stud. Oceanogr.* **2016**, *127*, 41-52.

(83) Williams, A.; MacKay, S., Differentiation of the brachiopod periostracum. *Palaeontology* **1979**, *22*, 721-736.

(84) Peck, L. S.; Clark, M. S.; Power, D.; Reis, J.; Batista, F. M.; Harper, E. M., Acidification effects on biofouling communities: winners and losers. *Global Change Biol.* **2015**, *21*, 1907-1913.

(85) Mackenzie, C. L.; Ormondroyd, G. A.; Curling, S. F.; Ball, R. J.; Whiteley, N. M.; Malham, S. K., Ocean warming, more than acidification, reduces shell strength in a commercial shellfish species during food limitation. *PLoS ONE* **2014**, *9*, e86764.

(86) Melzner, F.; Stange, P.; Trubenbach, K.; Thomsen, J.; Casties, I.; Panknin, U.; Gorb, S. N.; Gutowska, M. A., Food supply and seawater pCO<sub>2</sub> impact calcification and internal shell dissolution in the blue mussel *Mytilus edulis*. *PLoS ONE* **2011**, *6*, e24223.

(87) Day, E. G.; Branch, G. M.; Viljoen, C., How costly is molluscan shell erosion? A comparison of two patellid limpets with contrasting shell structures. *J. Exp. Mar. Biol. Ecol.* **2000**, *243*, 185-208.

- (88) Hoffman, J. I.; Peck, L. S.; Hillyard, G.; Zieritz, A.; Clark, M. S., No evidence for genetic differentiation between Antarctic limpet *Nacella concinna* morphotypes. *Mar. Biol.* **2009**, *157*, 765-778.
- (89) McDonald, M. R.; McClintock, J. B.; Amsler, C. D.; Rittschof, D.; Angus, R. A.; Orihuela, B.; Lutostanski, K., Effects of ocean acidification over the life history of the barnacle *Amphibalanus amphitrite*. *Mar. Ecol. Prog. Ser.* **2009**, *385*, 179-187.
- (90) Ye, F.; Jurikova, H.; Angiolini, L.; Brand, U.; Crippa, G.; Henkel, D.; Laudien, J.; Hiebenthal, C.; Šmajgl, D., Variation in brachiopod microstructure and isotope geochemistry under low-pH-ocean acidification conditions. *Biogeosciences* **2019**, *16*, 617-642.
- (91) Griesshaber, E.; Schmahl, W. W.; Neuser, R.; Pettke, T.; Blum, M.; Mutterlose, J.; Brand, U., Crystallographic texture and microstructure of terebratulide brachiopod shell calcite: An optimized materials design with hierarchical architecture. *Am. Mineral.* **2007**, *92*, 722-734.
- (92) Pérez-Huerta, A.; Dauphin, Y.; Cusack, M., Biogenic calcite granules--are brachiopods different? *Micron* **2013**, *44*, 395-403.
- (93) Harper, E. M., Are calcitic layers an effective adaptation against shell dissolution in the Bivalvia? *J. Zool. Lond.* **2000**, *251*, 179-186.
- (94) Cross, E. L.; Harper, E. M.; Peck, L. S., A 120-year record of resilience to environmental change in brachiopods. *Global Change Biol.* **2018**, *24*, 2262-2271.
- (95) Palmer, A. R., Calcification in marine molluscs: How costly is it? *PNAS* **1992**, *89*, 1379-1382.
- (96) Spalding, C.; Finnegan, S.; Fischer, W. W., Energetic costs of calcification under ocean acidification. *Global Biogeochem. Cycles* **2017**.
- (97) Watson, S.-A.; Morley, S. A.; Peck, L. S., Latitudinal trends in shell production cost from the tropics to the poles. *Sci. Adv.* **2017**, *3*, e1701362.
- (98) Ghedini, G.; Connell, S. D., Organismal homeostasis buffers the effects of abiotic change on community dynamics. *Ecology* **2016**, *97*, 2671-2679.

830 (99) Connell, S. D.; Ghedini, G., Resisting regime-shifts: the stabilising effect of  
831 compensatory processes. *Trends Ecol. Evol.* **2015**, *30*, 513-5.

832

833 DISCLOSURES

834 The authors declare no competing financial interest.